Boundary conditions for Boussinesq-type models in elevation-flux form

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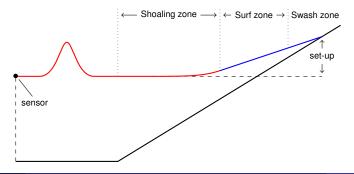


Objectives

Postdoc supported by **Institut des Mathématiques de la Planète Terre** Supervision : David Lannes and Philippe Bonneton

Long term goal: study extreme waves in littoral area

- Need accurate dispersive model: Boussinesq-type systems
- Boundary conditions are difficult to deal with Recently: Perfectly Matched Layer, source function method → costly

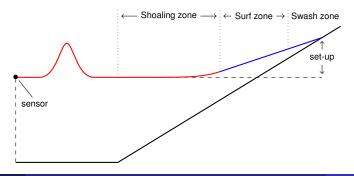


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- Need accurate dispersive model: Boussinesq-type systems
- Boundary conditions are difficult to deal with Recently: Perfectly Matched Layer, source function method → costly
- \rightarrow We propose a new and efficient method for boundary conditions.



Consider the Boussinesq-Abbott system

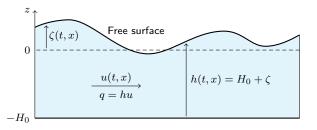
$$\partial_t \zeta + \partial_x q = 0$$
 in (0, ℓ) (BA)
(1 - $\kappa^2 \partial_{xx}^2) \partial_t q + \partial_x f_{\text{NSW}}(\zeta, q) = 0$

with generating boundary conditions

$$\zeta(t,0) = g_0(t), \qquad \zeta(t,\ell) = g_\ell(t),$$
 (1)

where $g_0, g_\ell \in C(0, T)$ and

$$\kappa^2 = H_0^2/3, \qquad f_{\rm NSW}(\zeta,q) = h u^2 + g h^2/2$$



How to account for boundary conditions? How to recover $q_{|_{x=0,\ell}}$?

- Hyperbolic case ($\kappa = 0$) : Riemann invariants
- Dispersive case ($\kappa > 0$) : need to invert $(1 \kappa^2 \partial_{xx}^2) \rightarrow$ requires knowledge on $\partial_t q_{|_{x=0,\ell}}$ Lannes and Weynans 2020 "Generating boundary conditions for a Boussinesq system"

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Fix $0 \le t \le T$, then $y(x) = \partial_t q(t, x)$ satisfies an ODE of the form $\begin{cases} y - \kappa^2 y'' = \phi(x) \\ y(0) = \dot{q}_{l_0}, \quad y(\ell) = \dot{q}_{l_\ell} \end{cases}$ Equivalently: $y = y_h + y_b$ with $\begin{cases} y_h - \kappa^2 y''_h = 0 \\ y_h(0) = \dot{q}_{l_0}, \quad y_h(\ell) = \dot{q}_{l_\ell} \end{cases}$ and $\begin{cases} y_b - \kappa^2 y''_b = \phi(x) \\ y_b(0) = y_b(\ell) = 0 \end{cases}$ How to account for boundary conditions? How to recover $q_{|_{x=0,\ell}}$?

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Define R^0 as the inverse of $(1 - \kappa^2 \partial_{xx}^2)$ with **homogeneous Dirichlet** conditions. Then

$$\partial_{t}q = \underbrace{-R^{0}\partial_{x}f_{\text{NSW}}}_{y_{b}} + \underbrace{\mathfrak{S}_{(0)}\dot{q}_{|_{0}} + \mathfrak{S}_{(\ell)}\dot{q}_{|_{\ell}}}_{y_{h}}$$
where
$$\begin{cases} (1 - \kappa^{2}\partial_{x}^{2})\mathfrak{S}_{(0)} = 0\\ \mathfrak{S}_{(0)}(0) = 1, \quad \mathfrak{S}_{(0)}(\ell) = 0 \end{cases} \text{ and } \begin{cases} (1 - \kappa^{2}\partial_{x}^{2})\mathfrak{S}_{(\ell)} = 0\\ \mathfrak{S}_{(\ell)}(0) = 0, \quad \mathfrak{S}_{(\ell)}(\ell) = 1 \end{cases} .$$
(2)

Reformulation of the model

Note R^1 the inverse of $(1 - \kappa^2 \partial_{xx}^2)$ with **homogeneous Neumann** conditions

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Proposition 1 (Equivalent formulation with nonlocal flux)

Let (ζ, q) initially equal to (ζ^{in}, q^{in}) . The two assertions are equivalent:

- The pair (ζ , q) satisfies (BA) with generating conditions $\zeta(\cdot, 0) = g_0$ and $\zeta(\cdot, \ell) = g_\ell$
- 2 The pair (ζ, q) satisfies

with

$$\begin{pmatrix} \mathfrak{s}_{(0)}'(0) & \mathfrak{s}_{(\ell)}'(0) \\ \mathfrak{s}_{(0)}'(\ell) & \mathfrak{s}_{(\ell)}'(\ell) \end{pmatrix} \begin{pmatrix} \dot{q}_0 \\ \dot{q}_\ell \end{pmatrix} = \frac{1}{\kappa^2} \begin{pmatrix} (R^1 - \mathrm{id})_{|_0} f_{\mathrm{NSW}} \\ (R^1 - \mathrm{id})_{|_\ell} f_{\mathrm{NSW}} \end{pmatrix} - \begin{pmatrix} \ddot{g}_0 \\ \ddot{g}_\ell \end{pmatrix}$$
(4)

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with

Sketch of the proof:

- To get (3), check that $R^0 \partial_x = \partial_x R^1$.
- Apply ∂_x to the discharge eq. from (3); take the trace at $x = 0, \ell$ to get (4).

$$\underbrace{\partial_{xt}^2 \boldsymbol{q}}_{-\partial_{tt}^2 \boldsymbol{\zeta}} + \underbrace{(\partial_{xx} R^1 f_{\text{NSW}})}_{\frac{1}{k^2} (\text{id} - R^1) f_{\text{NSW}}} = \mathfrak{s}'_{(0)} \dot{\boldsymbol{q}}_0 + \mathfrak{s}'_{(\ell)} \dot{\boldsymbol{q}}_{\ell}$$

Possibility to enforce general boundary conditions

$$\xi^{+}[\zeta, q](t, 0) = g_{0}(t), \qquad \xi^{-}[\zeta, q](t, \ell) = g_{\ell}(t).$$

For instance, ξ^{\pm} given by q or Riemann invariants

$$\mathcal{R}_{\pm}(U) = u \pm 2\sqrt{gh}.$$

(5)

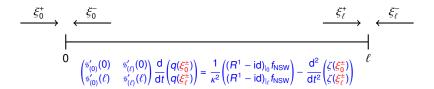
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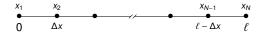
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Adapt trace ODE in terms of missing data (outgoing information ξ_0^- and ξ_ℓ^+)



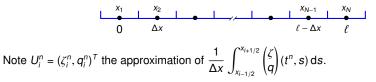
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Discretize $(0, \ell)$ as follows:



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Discretize $(0, \ell)$ as follows:

$$U_{i}^{n} = (\zeta_{i}^{n}, q_{i}^{n})^{T} \text{ the approximation of } \frac{1}{\Delta x} \int_{x_{i-1/2}}^{x_{i+1/2}} \left(\zeta_{q}^{i} \right) (t^{n}, s) \, \mathrm{d}s.$$

Time stepping procedure

Ν

Step 1: Define $\underline{R}^1 f_{\text{NSW}}^n$ as the vector $v \in \mathbb{R}^N$ satisfying

$$\begin{cases} v_{i} - \kappa^{2} \frac{v_{i+1} - 2v_{i} + v_{i-1}}{\Delta x^{2}} = f_{\text{NSW}}(U_{i}^{n}) \text{ for } 2 \le i \le N - 1\\ \frac{v_{2} - v_{1}}{\Delta x} = \frac{v_{N} - v_{N-1}}{\Delta x} = 0 \end{cases}$$

Similar definition for $\underline{\mathfrak{S}}_{(0)}$ and $\underline{\mathfrak{S}}_{(\ell)}$.

Time stepping procedure

Step 2: Approx. trace ODEs using FD scheme to get $\delta_t q_1^n, \delta_t q_N^n$; Update border values

$$\begin{cases} q_1^{n+1} = q_1^n + \Delta t \delta_t q_1^n \\ q_N^{n+1} = q_N^n + \Delta t \delta_t q_N^n \end{cases} \text{ and } \begin{cases} \zeta_1^{n+1} = g_0(t^{n+1}) \\ \zeta_N^{n+1} = g_\ell(t^{n+1}) \end{cases}$$

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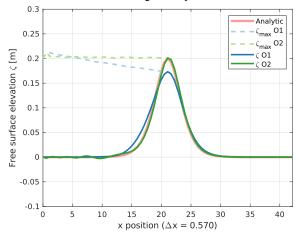
$$\begin{cases} q_1^{n+1} = q_1^n + \Delta t \delta_t q_1^n \\ q_N^{n+1} = q_N^n + \Delta t \delta_t q_N^n \end{cases} \text{ and } \begin{cases} \zeta_1^{n+1} = g_0(t^{n+1}) \\ \zeta_N^{n+1} = g_\ell(t^{n+1}) \end{cases}$$

Step 3: For $2 \le i \le N$, finite volume update with Lax-Friedrichs numerical flux

$$\begin{bmatrix} \frac{\zeta_i^{n+1} - \zeta_i^n}{\Delta t} + \frac{1}{\Delta x} \left(q_{i+1/2}^n - q_{i-1/2}^n \right) = 0 \\ \frac{q_i^{n+1} - q_i^n}{\Delta t} + \frac{1}{\Delta x} \left((\underline{R}^1 f_{\text{NSW}}^n)_{i+1/2} - (\underline{R}^1 f_{\text{NSW}}^n)_{i-1/2} \right) = (\mathfrak{s}_{(0)})_i \delta_t q_1^n + (\mathfrak{s}_{(\ell)})_i \delta_t q_N^n$$

Numerical scheme for the reformulated system

- Second order extension: MacCormack (prediction-correction)
- Advantage: no sponge layer required



Incoming solitary wave

Δx	Lax-Friedrichs		MacCormack	
	L ² -error	Order	L ² -error	Order
0.569662	0.052002	_	0.020162	_
0.284831	0.040773	0.35	0.003910	2.37
0.142416	0.024022	0.76	0.000767	2.35
0.071208	0.012777	0.91	0.000161	2.25
0.035604	0.006621	0.95	0.000037	2.12

Table: Error for incoming soliton (ζ enforced)

Δx	Lax-Friedrichs		MacCormack	
	L ² -error	Order	L ² -error	Order
0.284831	0.024246	_	0.001088	-
0.142416	0.014574	0.73	0.000314	1.79
0.071208	0.008471	0.78	0.000101	1.64
0.035604	0.004751	0.83	0.000029	1.80
0.017802	0.002559	0.89	0.000008	1.86

Table: Error for outgoing soliton (ζ enforced)

Account for varying bottom with Boussinesq-Peregrine in (ζ, q) -coordinates

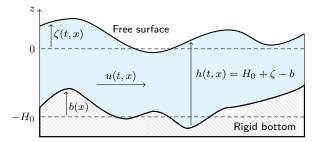
$$\begin{cases} \partial_t \zeta + \partial_x q = 0 \\ (1 + h_b \mathcal{T}_b) \partial_t q + \partial_x f_{\text{NSW}} = -gh\partial_x b \end{cases} \quad \text{in } (0, \ell) ,$$
 (BP)

under generating boundary conditions

$$\zeta(t,0)=g_0(t),\qquad \zeta(t,\ell)=g_\ell(t),$$

with $h_b = H_0 - b$ (depth at rest) and

$$\mathcal{T}_{b}(\cdot) = -\frac{1}{3h_{b}}\partial_{x}\left(h_{b}^{3}\partial_{x}\frac{(\cdot)}{h_{b}}\right) + \frac{(\cdot)}{2}\partial_{x}^{2}b,$$
(6)



Note R_b^0 the inverse of $(1 + h_b T_b)$ with **homogeneous Dirichlet** conditions. Then

$$\partial_t q = -R_b^0 \partial_x f_{\text{NSW}} - g R_b^0 (h \partial_x b) + \mathfrak{s}_{(b,0)} \dot{q}_{|_0} + \mathfrak{s}_{(b,\ell)} \dot{q}_{|_\ell}$$
(7)

$$\begin{array}{l} (1+h_b\mathcal{T}_b)\mathfrak{s}_{(b,0)} = 0 \\ \mathfrak{s}_{(b,0)}(0) = 1, \quad \mathfrak{s}_{(b,0)}(\ell) = 0 \end{array} \quad \text{and} \quad \begin{cases} (1+h_b\mathcal{T}_b)\mathfrak{s}_{(b,\ell)} = 0 \\ \mathfrak{s}_{(b,\ell)}(0) = 0, \quad \mathfrak{s}_{(b,\ell)}(\ell) = 1 \end{cases}$$

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where

$$\text{ere } \begin{cases} (1+h_b\mathcal{T}_b)\mathfrak{s}_{(b,0)} = 0\\ \mathfrak{s}_{(b,0)}(0) = 1, \quad \mathfrak{s}_{(b,0)}(\ell) = 0 \end{cases} \text{ and } \begin{cases} (1+h_b\mathcal{T}_b)\mathfrak{s}_{(b,\ell)} = 0\\ \mathfrak{s}_{(b,\ell)}(0) = 0, \quad \mathfrak{s}_{(b,\ell)}(\ell) = 1 \end{cases}$$

Lemma 1 (generalization of $R^0 \partial_x = \partial_x R^1$)

We can construct a nonlocal operator R_b^1 such that

$$\mathbf{R}_{b}^{0}\partial_{x}(\cdot) = \frac{1}{\alpha}(\partial_{x} + \phi)\left[h_{b}^{2}\mathbf{R}_{b}^{1}\left(\frac{(\cdot)}{h_{b}^{2}}\right)\right] - \mathbf{R}_{b}^{0}\left((\cdot)\phi\right) \quad \text{with} \quad \alpha = 1 + \frac{1}{4}(\partial_{x}b)^{2} \quad \text{and} \quad \phi = \frac{3}{2}\frac{\partial_{x}b}{h_{b}}$$

Definition 1 (Nonlocal flux and source terms)

$$\tilde{\mathfrak{f}} = h_b^2 R_b^1 \left(\frac{f_{\rm NSW}}{h_b^2} \right), \qquad \mathfrak{S} = R_b^0 (-gh\partial_x b) - \frac{\phi}{\alpha} \tilde{\mathfrak{f}} + R_b^0 (\phi f_{\rm NSW})$$

Proposition 2 (Equivalent formulation with nonlocal flux)

Let (ζ, q) initially equal to (ζ^{in}, q^{in}) . The two assertions are equivalent:

- The pair (ζ, q) satisfies (BP) with generating conditions $\zeta(\cdot, 0) = g_0$ and $\zeta(\cdot, \ell) = g_\ell$
- 2 The pair (ζ, q) satisfies

$$\partial_{t}\zeta + \partial_{x}q = 0,$$

$$\partial_{t}q + \frac{1}{\alpha}\partial_{x}\mathfrak{f}(U, x) = \mathfrak{S}(U, x) + \mathfrak{S}_{(b,0)}\dot{q}_{0} + \mathfrak{S}_{(b,\ell)}\dot{q}_{\ell}$$
in (0, ℓ) (8)

and the trace equations

$$\begin{pmatrix} s'_{(b,0)}(0) & s'_{(b,\ell)}(0) \\ s'_{(b,0)}(\ell) & s'_{(b,\ell)}(\ell) \end{pmatrix} \begin{pmatrix} \dot{q}_0 \\ \dot{q}_\ell \end{pmatrix} = \Phi \Big(q_{|_{0,\ell}}, \hat{\eta}_{|_{0,\ell}}, \partial_x R^0_{b|_{0,\ell}}(\phi f_{\mathsf{NSW}} - gh\partial_x b) \Big) - \begin{pmatrix} \ddot{g}_0 \\ \ddot{g}_\ell \end{pmatrix}$$
(9)

where $\Phi : \mathbb{R}^3 \to \mathbb{R}^2$ known.

Question: starting from a wrong initial condition, can we recover the solution by enforcing appropriate boundary conditions?

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Setup:

- approximate U_{ref} solution in $(-\ell, 2\ell)$ with periodic conditions;
- extract $g_0(t^n) := \sigma(t^n)\xi^+[U_{\text{ref}}]_{|_0}(t^n)$ and $g_\ell(t^n) := \sigma(t^n)\xi^-[U_{\text{ref}}]_{|_\ell}(t^n)$;
- approximate new solution in $(0, \ell)$, initially at rest, with g_0, g_ℓ enforced at boundaries;

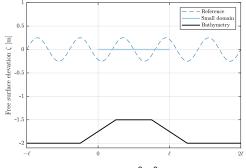
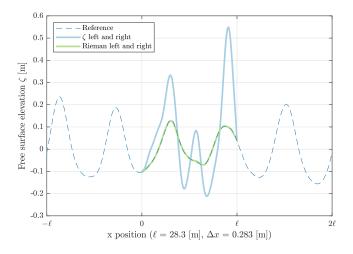
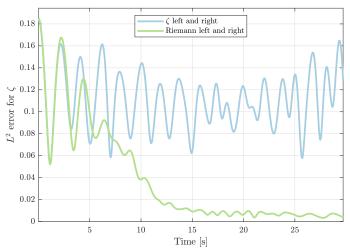


Figure: Initial condition ($(2\pi H_0)^2/\lambda^2 = 0.31, a_0/H_0 = 0.25$)

Comparing different boundary conditions

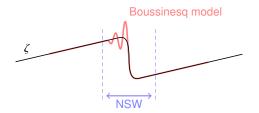




Error to reference solution for different boundary conditions

Coupling Boussinesq-Peregrine and shallow water models

Motivation: wave breaking with dispersive models \rightarrow non physical oscillations.



 \rightarrow Cancel dispersive term near shock wave

$$\begin{cases} \partial_{t}\zeta_{L} + \partial_{x}q_{L} = 0 & \text{in } (0, \ell_{1}) \\ \partial_{t}q_{L} + \frac{1}{\alpha}\partial_{x}f(U_{L}) = \mathfrak{S}(U_{L}) + \mathfrak{s}_{(b,0)}\dot{q}_{L|_{0}} + \mathfrak{s}_{(b,\ell_{1})}\dot{q}_{L|_{\ell_{1}}} & \text{in } (0, \ell_{1}) \\ \partial_{t}\zeta_{R} + \partial_{x}q_{R} = 0 & \text{in } (\ell_{1}, \ell_{2}) \\ \partial_{t}q_{R} + \partial_{x}f_{NSW}(U_{R}) = -gh_{R}\partial_{x}b & \text{in } (\ell_{1}, \ell_{2}) \end{cases}$$

Coupling conditions: $\mathcal{R}_+(U_R)_{|_{\ell_1}} = \mathcal{R}_+(U_L)_{|_{\ell_1}}, \quad \mathcal{R}_-(U_L)_{|_{\ell_1}} = \mathcal{R}_-(U_R)_{|_{\ell_1}}$



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In practice, overlapping helps to reduce oscillations/reflections

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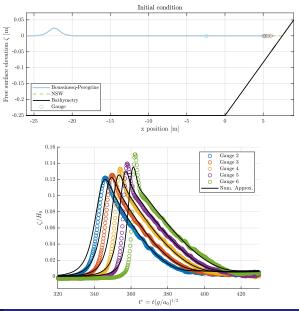
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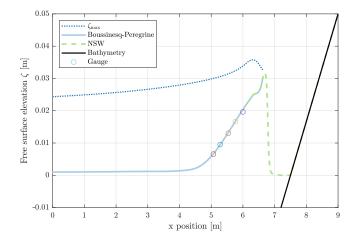
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Approx. U_Rⁿ⁺¹ with FV scheme + hydrostatic reconstruction; R₊(U_R)_{|ℓ1} = R₊(U_L)_{|ℓ1}
 Approx. U_Lⁿ⁺¹ with Lax-Friedrichs scheme + trace equations; R₋(U_L)_{|ℓ1+ε} = R₋(U_R)_{|ℓ1+ε}
 Convex combination in overlapping area: U_iⁿ⁺¹ = ρ(x_i)U_{Li}ⁿ⁺¹ + (1 - ρ(x_i))U_{Ri}ⁿ⁺¹.

Experimental testcase: LEGI



Mathieu Rigal



Approximate transparent boundary conditions

Use coupling as a sponge layer to evacuate waves.

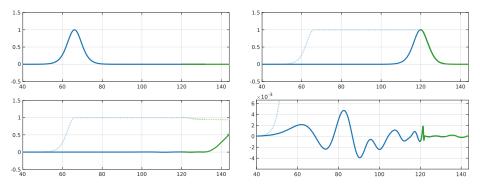


Figure: Outgoing soliton at times t = 0, 9.46, 14.19, and 23.16 [s]. Green domain corresponds to NSW.

Conclusion and perspectives

Over a flat bottom:

- Reformulation of Boussinesq-Abbott
- Generalized boundary conditions
- Efficient 1st and 2nd order schemes

Over a varying bottom:

- Approach extended to Boussinesq-Peregrine
- Coupling with shallow water model
- Implementation + validation (experimental data, various boundary conditions tested)

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- Extension to Boussinesq models with improved dispersion relation
- Statistics of extreme waves: impact of bathymetry

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Thank you for your attention!